

H Ch 11

FIND EQUIVALENCE POINT FIRST

CORRECT MOLARITY AS TITRANT IS ADDED

H 11-4 Polyprotic Titrations

homework for week 14,15 due dates this Wednesday and Friday

Titrations_

H Ch 11

Titration of a Strong Base with a Strong Acid

EX: 50.00 mL of 0.02000 M KOH titrated with 0.1000 M HBr.

chemical equation (why reaction arrow?) KOH(aq) + HBr(aq) -> KBr(aq) + H₂O(l) net ionic equation

 $H^+(aq) + OH^-(aq) -> H_2O(l)$

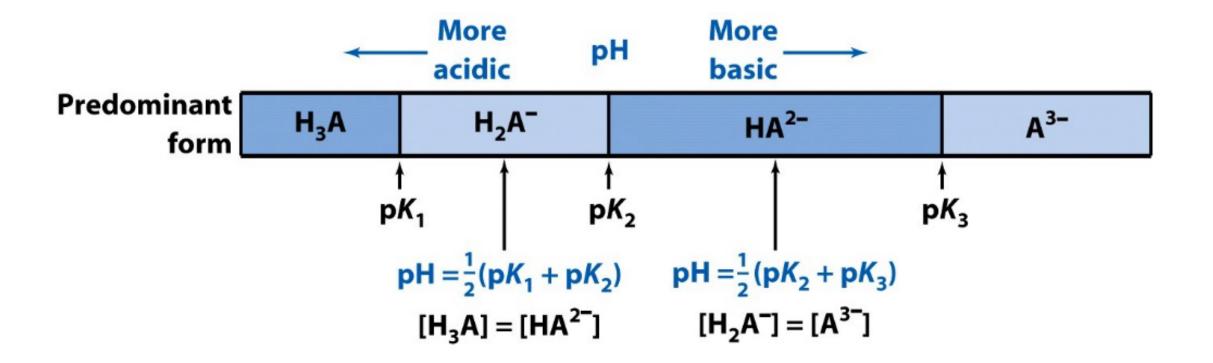
Titration of a Weak Acid with a Strong Base

EX: 50.00 mL of 0.02000 M MES, $pK_a = 6.27$, titrated with 0.1000 M NaOH.

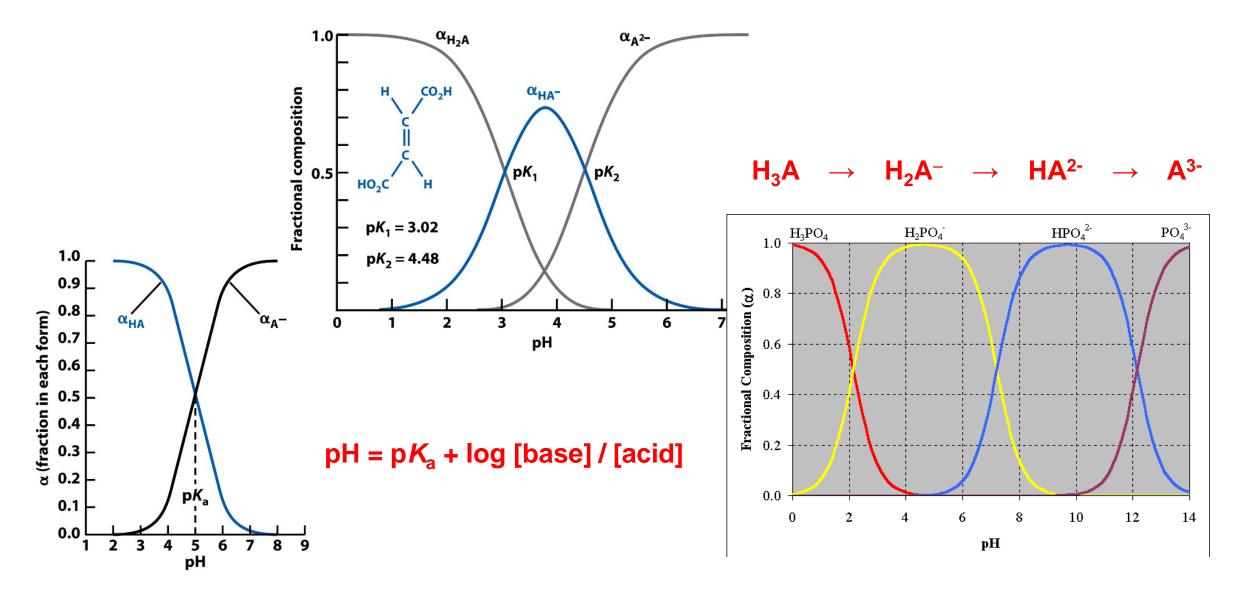
chemical equation (why reaction arrow?)net ionic equationNaOH(aq) + HA(aq) -> NaA(aq) + H2O(l)HA(aq) + OH^-(aq) -> A^-(aq) + H2O(l)

Finding the Principal Species at a Given pH

 $\mathbf{pH} = \mathbf{pK}_{a} + \log [base] / [acid]$



Finding the Fraction Dissociated at a Given pH

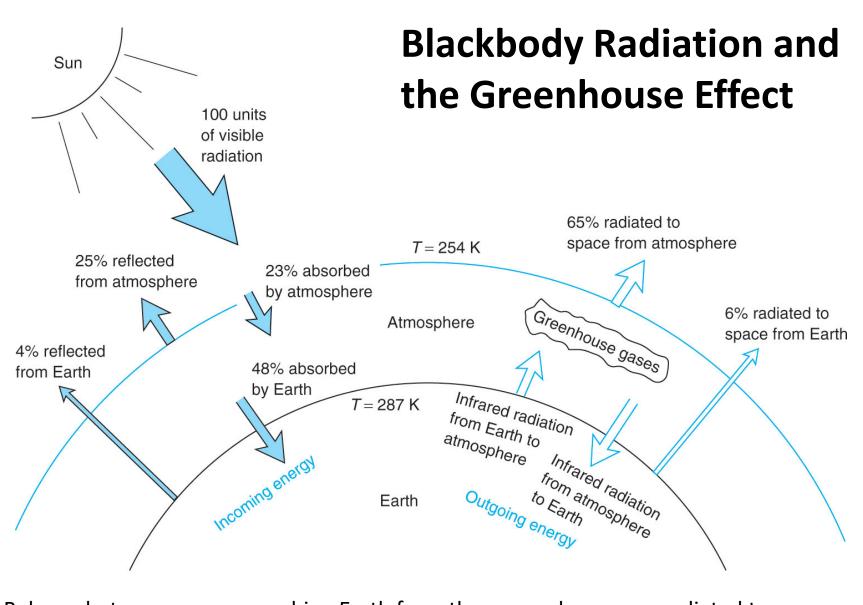


Leveling Effect: $2 H_2O(I) \ll H_3O^+(aq) + OH^-(aq)$

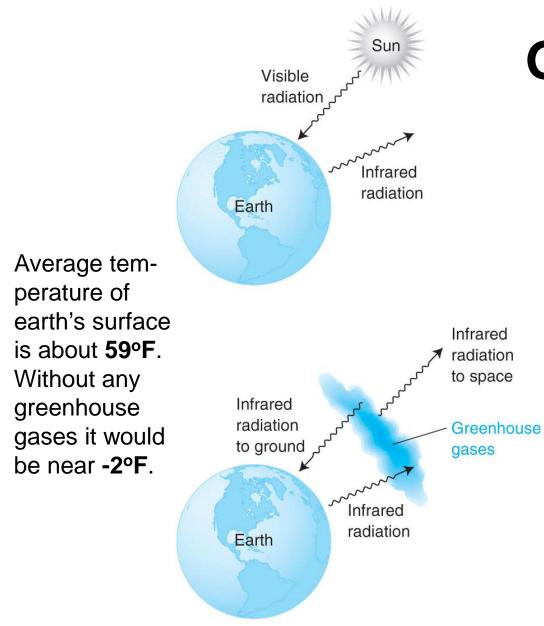
Strongest acid (base) that can exist in a solvent is the acidic (basic) autoionization species of the solvent.

Acidity Constants in Water at 25°C					Acidity Constants in Water at 25°C				
Acid	Formula	Conjugate Base	K,	pK,	Acid	Formula	Conjugate Base	K,	pK,
Hydriodic	HI	1-	≈ 10 ¹¹	≈ -11	Hydrazoic	HN ₃	N ₃	1.9×10^{-5}	4.72
Hydrobromic	HBr	Br	≈ 10 ⁹	≈ -9	Acetic	CH ₃ COOH	CH ₃ COO ⁻	1.8×10^{-5}	4.74
Perchloric	HCIO ₄	ClO ₄	≈ 10 ⁷	≈ -7	Propionic	CH3CH2COOH	CH ₃ CH ₂ COO ⁻	1.3×10^{-5}	4.89
Hydrochloric	HCI	CI-	≈ 10 ⁷	≈ -7	Pyridinium ion	HC ₅ H ₅ N ⁺	C ₃ H ₅ N (pyridine)	5.6×10^{-6}	5.25
Chloric	HCIO ₃	ClO ₃	≈ 10 ³	<i>≈</i> -3	Carbonic (1)	H ₂ CO ₃	HCO ₃	4.3×10^{-7}	6.37
Sulfuric (1)	H ₂ SO ₄	HSO ₄	$\approx 10^2$	≈ -2	Sulfurous (2)	HSO ₃	SO ₃ ²⁻	1.0×10^{-7}	7.00
Nitric	HNO3	NO ₃	≈ 20	≈ -1.3	Arsenic (2)	H ₂ AsO ₄ ⁻	HAsO ₄ ²⁻	9.3×10^{-8}	7.03
Hydronium ion	H_3O^+	H ₂ O	1	0.0	Hydrosulfuric	H ₂ S	HS ⁻	9.1×10^{-8}	7.04
Urea acidium ion	(NH ₂)CONH ⁺ ₃	(NH ₂) ₂ CO (urea)	$6.6 imes 10^{-1}$	0.18	Phosphoric (2)	H ₂ PO ₄ ⁻	HPO ₄ ²⁻	6.2×10^{-8}	7.21
Iodic	HIO ₃	IO ₃	1.6×10^{-1}	0.80	Hypochlorous	HCIO	CIO-	3.0×10^{-8}	7.52
Oxalic (1)	$H_2C_2O_4$	$HC_2O_4^-$	5.9×10^{-2}	1.23	Hydrocyanic	HCN	CN ⁻	6.2×10^{-10}	9.21
Sulfurous (1)	H ₂ SO ₃	HSO ₃	1.5×10^{-2}	1.82	Ammonium ion	NH ⁺	NH ₃	5.6×10^{-10}	9.25
Sulfuric (2)	HSO_4^-	SO_4^{2-}	1.2×10^{-2}	1.92	Carbonic (2)	HCO ₃	CO3 ²⁻	4.8×10^{-11}	10.32
Chlorous	HClO ₂	ClO ₂	1.1×10^{-2}	1.96	Methylammonium ion	CH ₃ NH ⁺	CH ₃ NH ₂	2.3×10^{-11}	10.64
Phosphoric (1)	H ₃ PO ₄	$H_2PO_4^-$	7.5×10^{-3}	2.12	Arsenic (3)	HAsO ₄ ²⁻	AsO ₄ ³⁻	3.0×10^{-12}	11.52
Arsenic (1)	H ₃ AsO ₄	H ₂ AsO ₄	5.0×10^{-3}	2.30	Hydrogen peroxide	H_2O_2	HO_2^-	2.4×10^{-12}	11.62
Chloroacetic	CICH ₂ COOH	CICH ₂ COO ⁻	1.4×10^{-3}	2.85	Phosphoric (3)	HPO_4^{2-}	PO4 ³⁻	2.2×10^{-13}	12.66
Hydrofluoric	HF	F ⁻	$6.6 imes 10^{-4}$	3.18	Water	H-O	OH-	1.0×10^{-14}	14.00
Nitrous	HNO ₂	NO_2^-	4.6×10^{-4}	3.34	Hydrogen sulfide ion	HS-	S2-	1.0×10^{-19}	19.00
Formic	HCOOH	HCOO-	$1.8 imes 10^{-4}$	3.74	Hydrogen	H ₂	H-	1.0×10^{-33}	33.00
Benzoic	C ₆ H ₅ COOH	C ₆ H ₃ COO ⁻	6.5×10^{-5}	4.19	Ammonia	NH ₃	NH ₂	1.0×10^{-38}	38.00
Oxalic (2)	$HC_2O_4^-$	$C_2O_4^{2-}$	6.4×10^{-5}	4.19	Hydroxide ion	OH-	O ²⁺		

conjugate bases stronger than OH-



Balance between energy reaching Earth from the sun and energy reradiated to space.

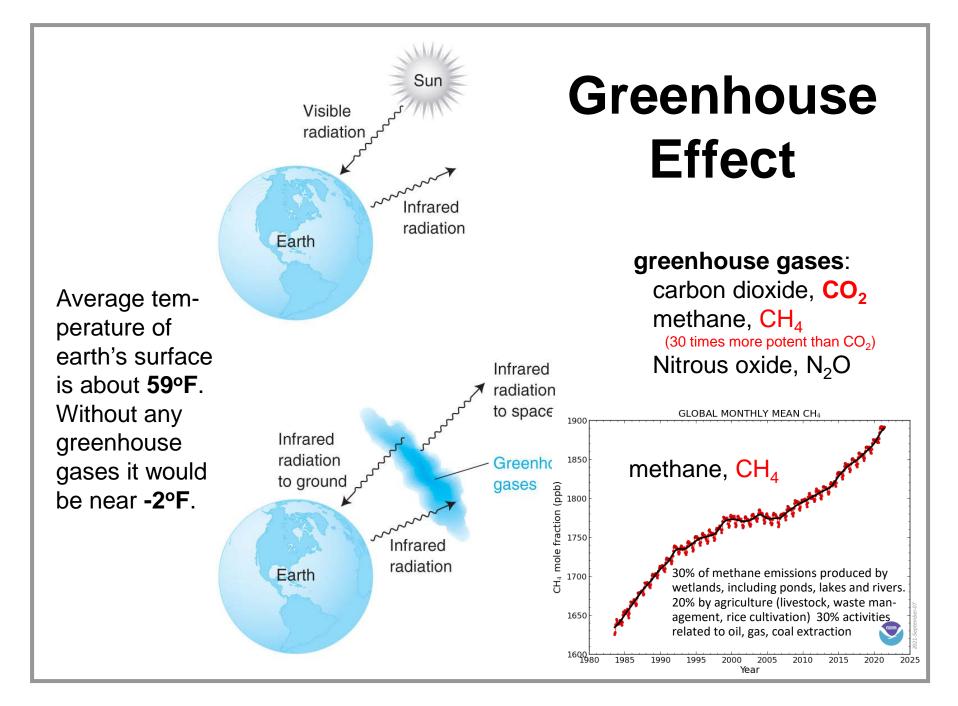


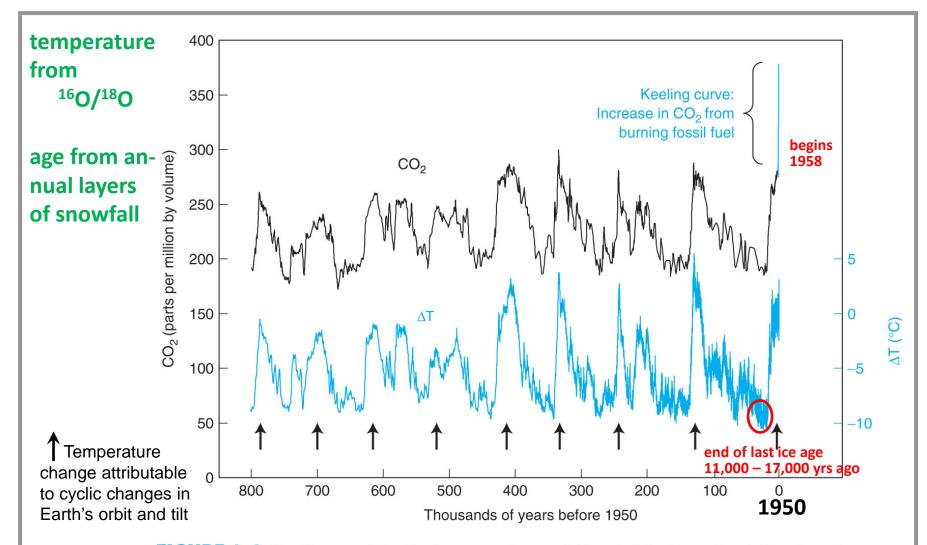
Greenhouse Effect

greenhouse gases: carbon dioxide, CO₂ methane, CH₄ (30 times more potent than CO₂) Nitrous oxide, N₂O

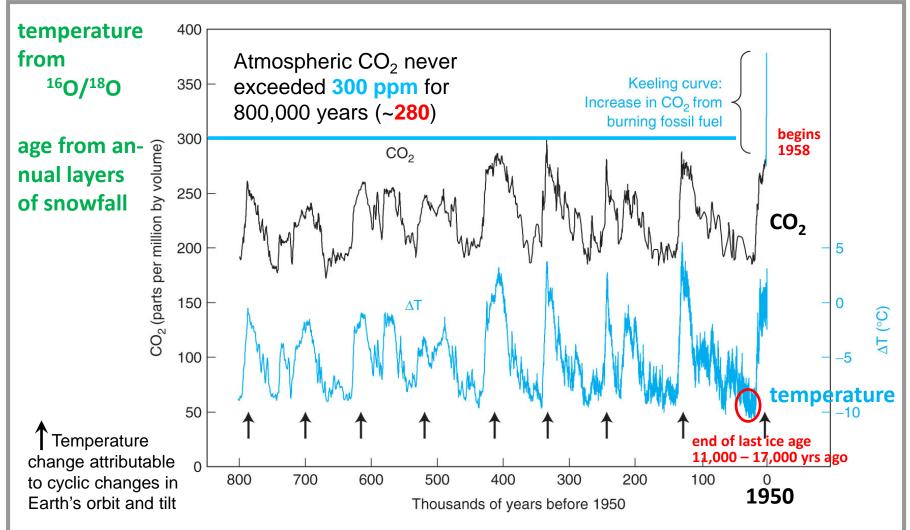
FIGURE 0-5 Greenhouse effect. The sun warms the Earth mainly with visible radiation. Earth emits infrared radiation, which would all go into space in the absence of the atmosphere. Greenhouse gases in the atmosphere absorb some of the infrared radiation and emit some of that radiation back to the Earth. Radiation directed back to Earth by greenhouse gases keeps the Earth warmer than it would be in the absence of greenhouse gases.

Harris, Quantitative Chemical Analysis, 8e





1998 - Russia, US, France in Antartica – 3623 meters! **FIGURE 0-6** Significance of the Keeling curve (upper right, color) is shown by plotting it on the same graph with atmospheric CO₂ measured in air bubbles trapped in ice cores drilled from Antarctica. Atmospheric temperature at the level where precipitation forms is deduced from hydrogen and oxygen isotopic composition of the ice. [Vostok ice core data from J. M. Barnola, D. Raynaud, C. Lorius, and N. I. Barkov, http://cdiac.esd.ornl.gov/ftp/trends/co2/vostok.icecore.co2.]



1998 - Russia, US, France in Antartica – 3623 meters! last time the atmospheric CO₂ amounts were this high was more than 3 million years ago, when temperature was 2°–3°C (**3.6°–5.4°F**) higher than during the pre-industrial era, and sea level was 15–25 meters (**50–80** feet) higher than today

Increase in CO₂ Over Past 60 Years

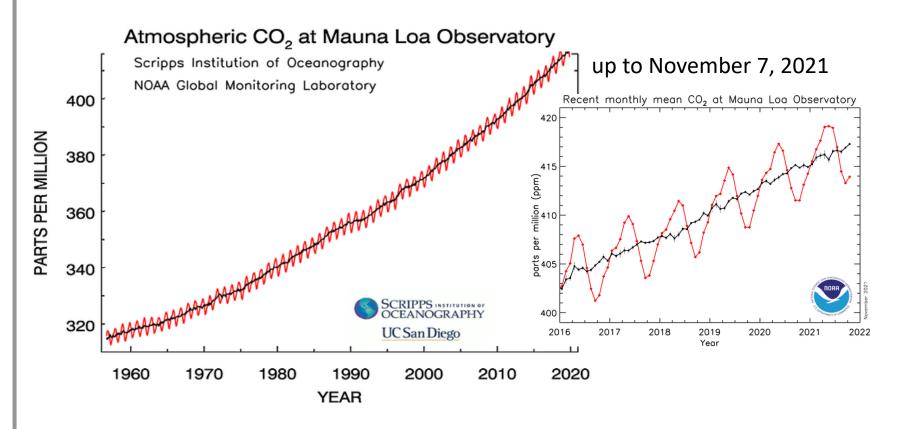
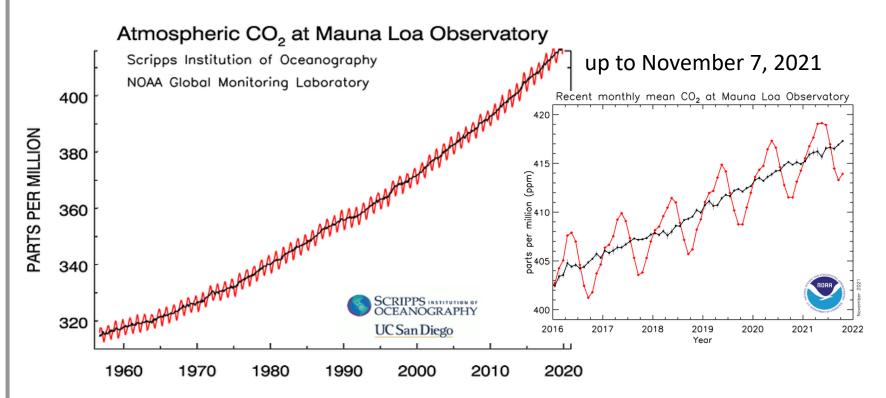


FIGURE 0-4 Monthly average atmospheric CO₂ measured on Mauna Loa. This graph, known as the *Keeling curve*, shows seasonal oscillations superimposed on rising CO₂. [Data from http://scrippsco2.ucsd.edu/data/ in_situ_co2/monthly_mlo.csv.]

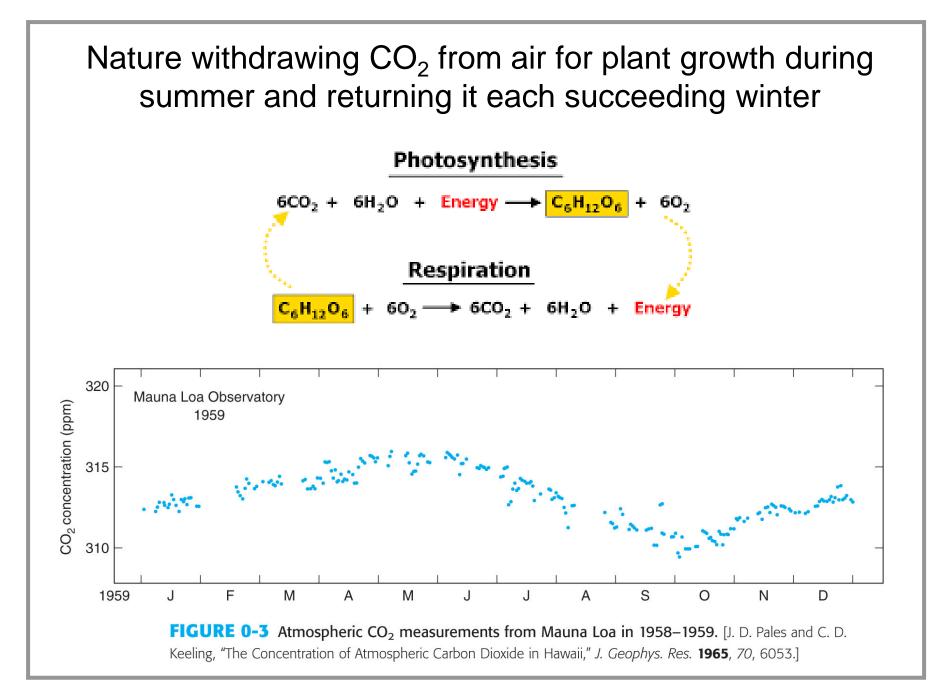
Historic CO_2 high of 280 ppm over 800,000 years increased to current 416 ppm in past 200 years due to burning fossil fuel (oil, coal, wood, natural gas) and destruction - "clearing" - of earth's forests.

Increase in CO₂ Over Past 60 Years

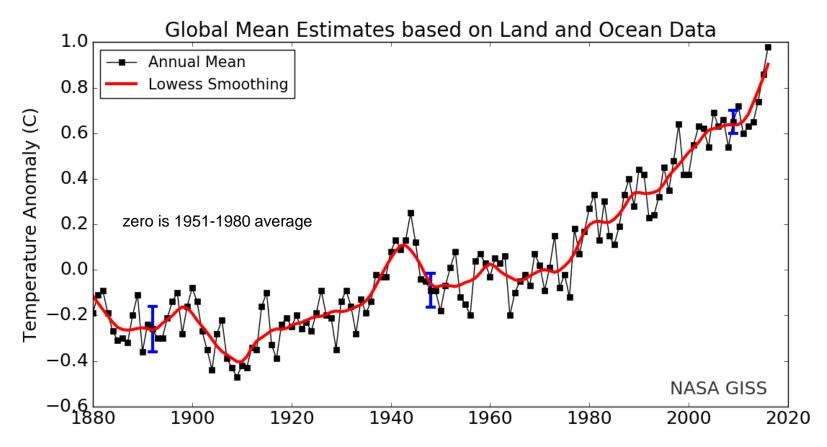


The atmospheric burden of CO_2 is now comparable to where it was during the Pliocene Climatic Optimum, between 4.1 and 4.5 million years ago, when CO_2 was close to, or above 400 ppm. During that time, sea level was about 78 feet higher than today, the average temperature was 7 degrees Fahrenheit higher than in pre-industrial times.

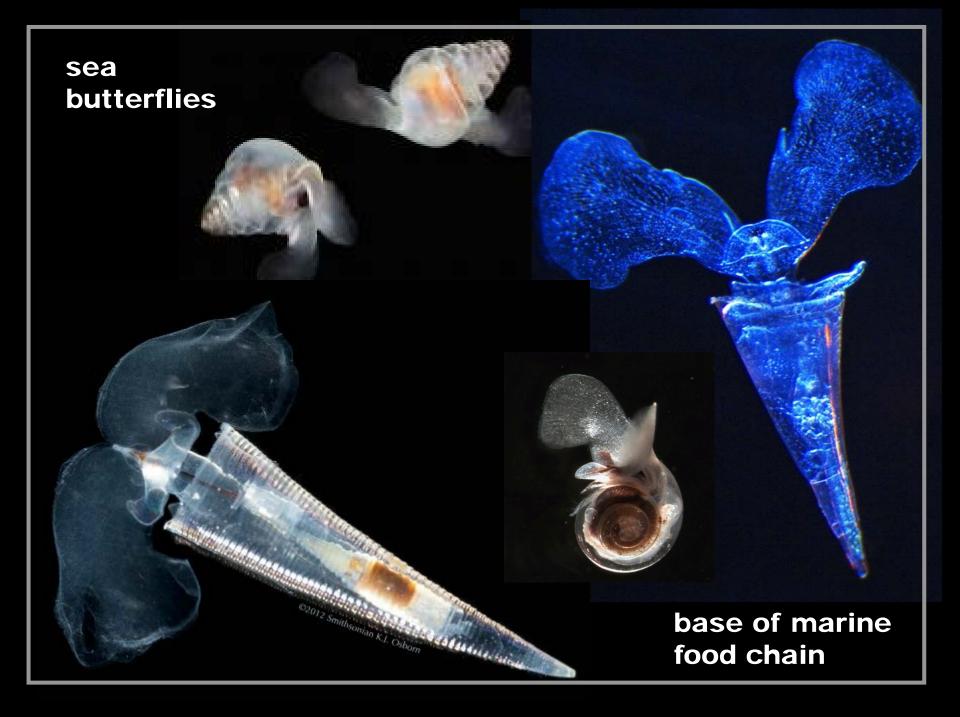
In February, the United States officially rejoined the Paris Agreement on climate change. Yet, as the measurements from Mauna Loa show, despite decades of negotiation, the global community has been unable to meaningfully slow, let alone reverse, annual increases in atmospheric CO₂ levels.



Global Warming



One effect of increasing levels of CO₂ in our atmosphere



Pteropod

Low carbonate concentration promotes dissolution of $CaCO_3$ shells and skeletons

 $CaCO_3(s) \rightleftharpoons Ca^{2+}(aq) + CO_3^{2-}(aq)$

Increasing atmospheric CO_2 increases concentration of CO_2 dissolved in oceans, which consumes carbonate and lowers the pH

$$\begin{array}{l} \mathrm{CO}_2(aq) + \mathrm{H_2O}(l) + \mathrm{CO_3^{2-}}(aq) \ \rightleftharpoons \ 2 \ \mathrm{HCO_3^{-}}(aq) \\ (\mathrm{H_2CO_3}) \end{array}$$

 H_2CO_3 : $K_{a1} = 4.5 \times 10^{-7}$, $K_{a2} = 4.7 \times 10^{-11}$

pre-industrial pH of oceans = 8.16 currently ocean pH = 8.04 without change, pH by 2100 = 7.7-7.8

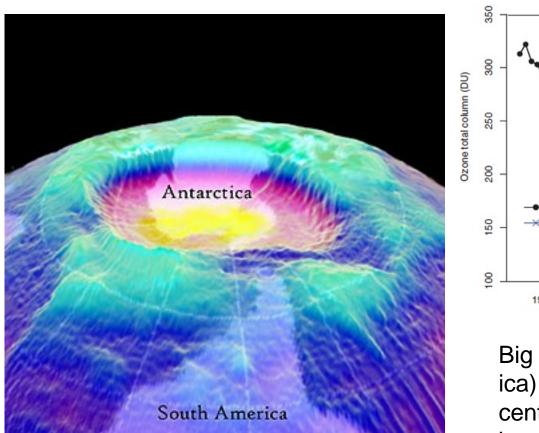
(earth not experienced such low pH in > 20 million years)

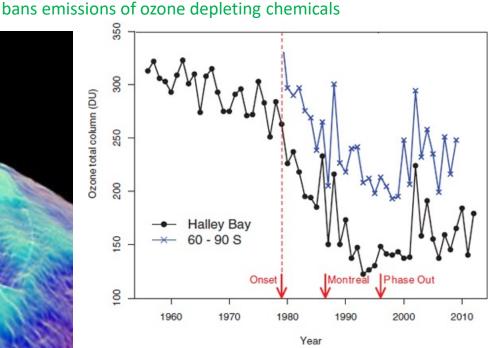


shell beginning to dissolve

The Ozone Hole

A success story – use of Freons stopped (Montreal Protocol) and hole is healing





Big hole (size of North America) still there but ozone concentration does not appear to be decreasing further.

(estimates than hole will "heal" – attain pre-1980 levels – by 2060)

FINAL Review

Final Exam: Tuesday, December 7, 8:00 – 10:00 am, 312 Lincoln Hall

On OWL ChemWorks Review Topics Mastery Review Topics On Website: Harris Review Topics Zumdahl Review Topics Equation Sheet for Final

FINAL Review

KNOWN TOPICS

Isotopes Lewis Structures Balancing Redox Reaction by Half-Reaction Method Titrations Nomenclature Stoichiometry Calculations Colligative Properties